

# Effect of Sub-pixel Clouds and Aerosols on the retrieval of Trace Gases

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# Effect of Clouds on the Retrieval of BL Trace Gas column

$$\hat{N} = \frac{SCD}{AMF}; \quad AMF = (1 - w) * AMF_{clear} + w * AMF_{cloud}$$

where,  $w$  is the "cloud fraction"

For clouds above polluted BL:  $AMF_{cloud} \approx 0$

$$\hat{N} \approx \frac{SCD}{(1 - w) AMF_{clear}}$$

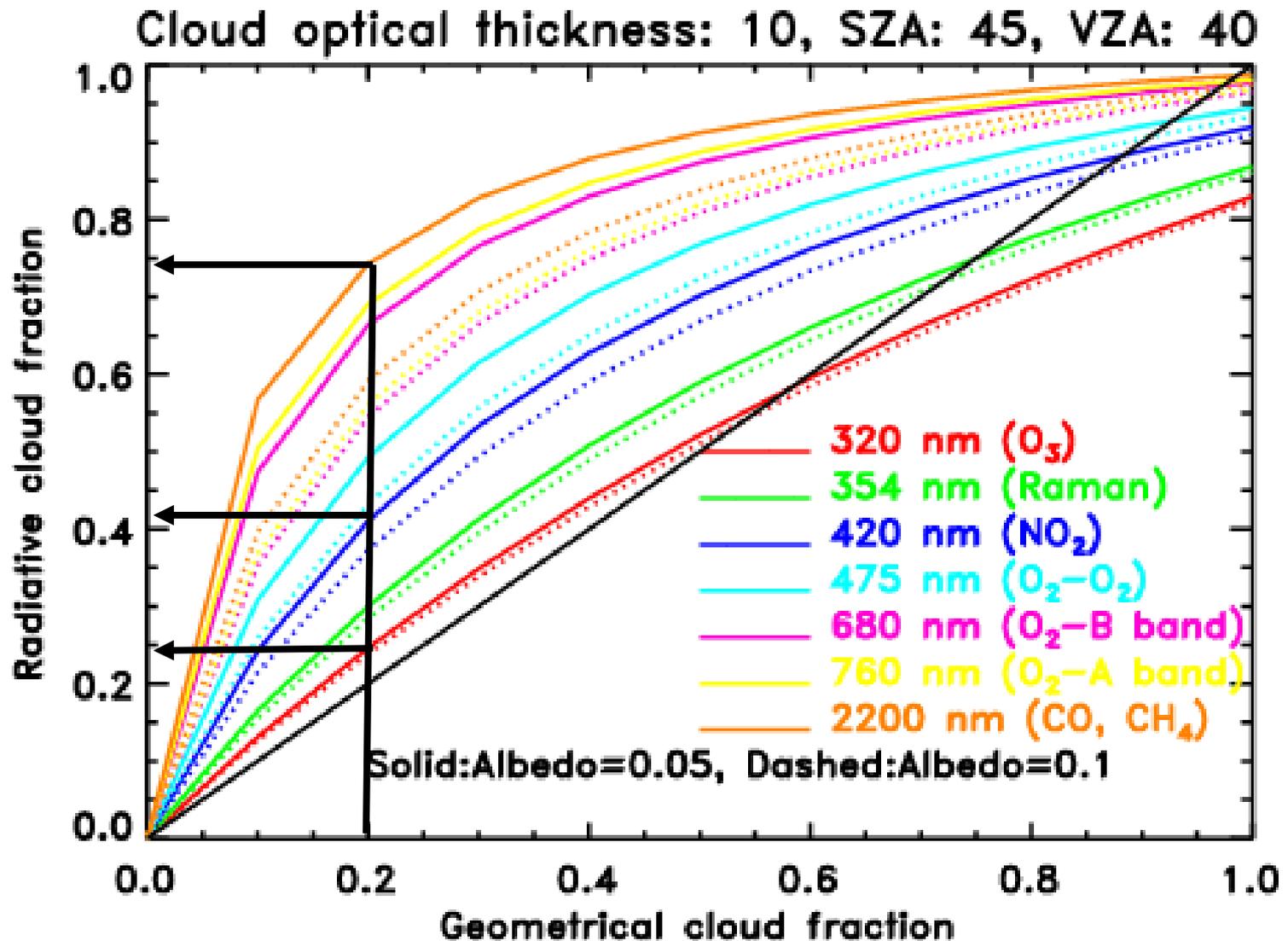
- Since clouds are usually much brighter than the surrounding atmosphere  $w$  can be several times larger than the geometrical cloud fraction.
- 10% error in  $w$  can produce 10-100% error in the estimated vert. column depending upon  $w$ .
- Cloud height is less important unless clouds are inside the BL.

# What is “Cloud Fraction”?

- For trace gas measurement at reflected  $\lambda$ s it is the fraction of the measured radiation that comes from clouds (not including below cloud radiance that propagates thru clouds).
  - The OMI team calls it Cloud Radiance Fraction or Radiative Cloud Fraction ( $w$ ).
- Formally,  $w = f_g * \rho_c / \rho_m$   
where  $f_g$  is the geometrical cloud fraction,  $\rho_c$  is the reflectance of the “bare” cloud, and  $\rho_m$  is the TOA reflectance corrected for atm absorption.

Cloud Radiance Fraction varies with  $\lambda$

# Rayleigh scattering effect on w



# A Simple Method to Calculate $w$

Using Mie theory it can be shown that  $w$  varies linearly with  $\rho_{clear} / \rho_m$

Applying boundary conditions :

$$w = \left( \frac{\rho_m - \rho_{clear}}{\rho_m^* - \rho_{clear}} \right) \left( \frac{\rho_m^*}{\rho_m} \right)$$

where,  $\rho_m^*$  is the "critical" TOA reflectance above which  $w = 1$

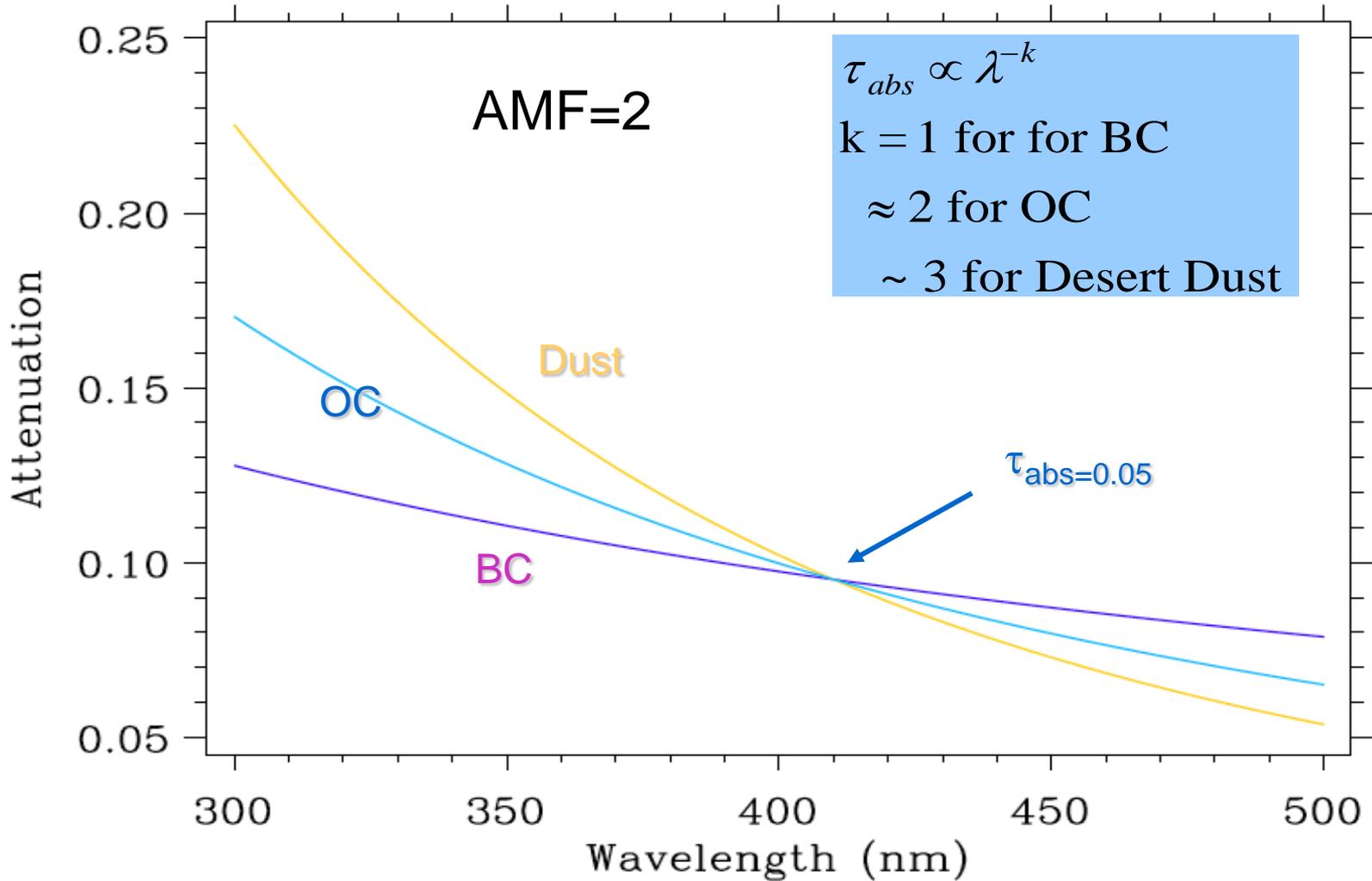
Since  $\rho_m^* \gg \rho_{clear}$ ,  $w$  is insensitive to  $\rho_m^*$

In the TOMS and OMI algorithms  $\rho_m^*$  is calculated assuming an 80% reflecting Lambertian cloud. The results agree well with Mie theory.

# Correction for Aerosols

- OMI method of calculating  $w$  accounts for non-absorbing aerosols of any kind.
- If aerosols have absorption,  $\rho_{\text{clear}}$  is overestimated, causing  $w$  to be underestimated.
  - the effect is worse in the UV than in the visible
  - the effect is worse for elevated aerosols
  - the effect is worse over brighter surfaces
- UV absorbing aerosol index (UV-AAI) can be used to develop a correction.

# Reduction in w by Aerosol Absorption

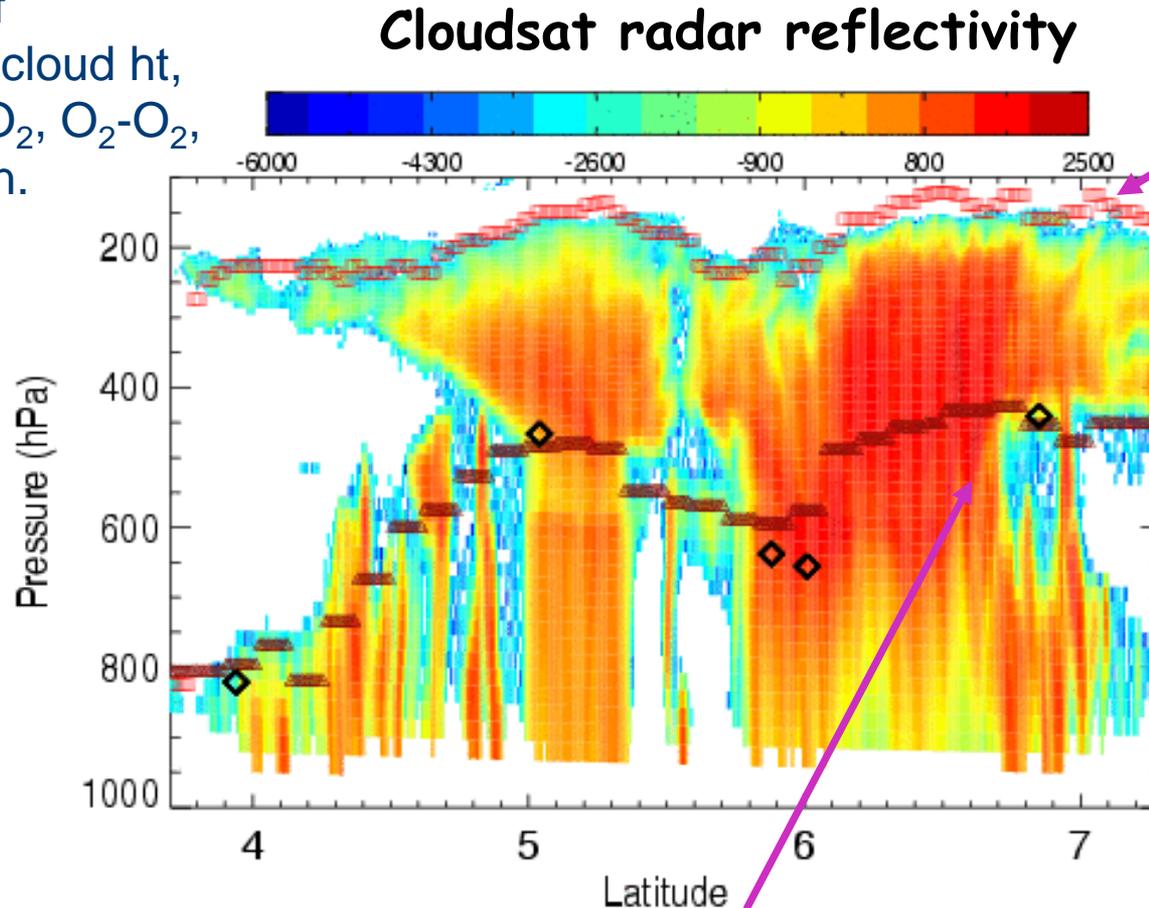


# What About Cloud height?

- At reflected wavelengths sunlight penetrates into clouds and passes through most types of clouds.
  - Hence trace gas absorption is affected by cloud vertical structure. Cloud/aerosol mixed scenes can produce complicated effects.
- Concept of Optical Centroid Pressure (OCP) has been developed to account for this complexity.
  - The concept is accurate only if the trace gas is well mixed between the scattering layers, but is ALWAYS more accurate than assuming cloud-top pressure to estimate AMF.

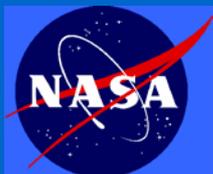
# Multi-phase/Multi-layer Cloud Effects

There are several methods of estimating cloud ht, e.g., TIR,  $O_2$ ,  $O_2-O_2$ , and Raman.



MODIS cloud-top press is insensitive to cloud vertical structure

Cloud Optical Centroid press calculated using OMI-measured Rot Raman Scattering is sensitive to cloud vert structure  
(ref : Vasilkov *et al.*, JGR, '08)



# Cloud OCP Retrieval Methods

- OMI uses Rotational Raman scattering @350 nm and O<sub>2</sub>-O<sub>2</sub> absorption @477 nm to estimate OCP.
  - Both methods agree well with Cloudsat-derived OCP when  $w > 0.5$ .
  - Random and systematic errors in OCP increase as  $1/w$  for small  $w$ .
  - Both methods overestimate OCP when  $w$  is small.

Retrieved BL trace gas columns are probably most accurate when  $0.3 < w < 0.7$

# Summary

- BL trace gas retrieval requires accurate knowledge of Cloud Radiance Fraction ( $w$ ).
  - A simple method to calculate  $w$ , developed independently at NASA and KNMI, works very well, but correction for absorbing aerosols needs to be developed.
- To estimate the AMF it is far better to use OCP than the cloud-top pressure.
  - The two OMI methods for estimating OCP underestimate the OCP of clouds and aerosols for  $w < 0.5$ .
- Aerosol and cloud mixed scenes present a significant challenge.
  - Currently, there is no good method to handle them.